

Precipitation and its variability in the high elevation area of the Nepal Himalayas

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Outline

- **HiPRECS:**
In-situ observation project in the eastern Nepal Himalayas
- **Diurnal cycle** of precipitation:
by land surface effects
in higher elevations by the glacier
- **Heavy rainfall event:**
by monsoon lows under typical ISO's influences
- Summary



Tsho Rolpa glacier lake and Trakarding Glacier in Rolwaling, eastern Nepal Himalayas ©Dr. Koji Fujita@NU

HiPRECS (Himalaya Precipitation Study)

Hydrological cycle in the Himalayas:

- **Large amounts of summer precipitation** in the central-eastern Himalayas
- **Glaciers:** summer accumulation type
- **Head waters of major rivers and water resources**

Precipitation and its variability in summer:

- essential element, but still poorly understood

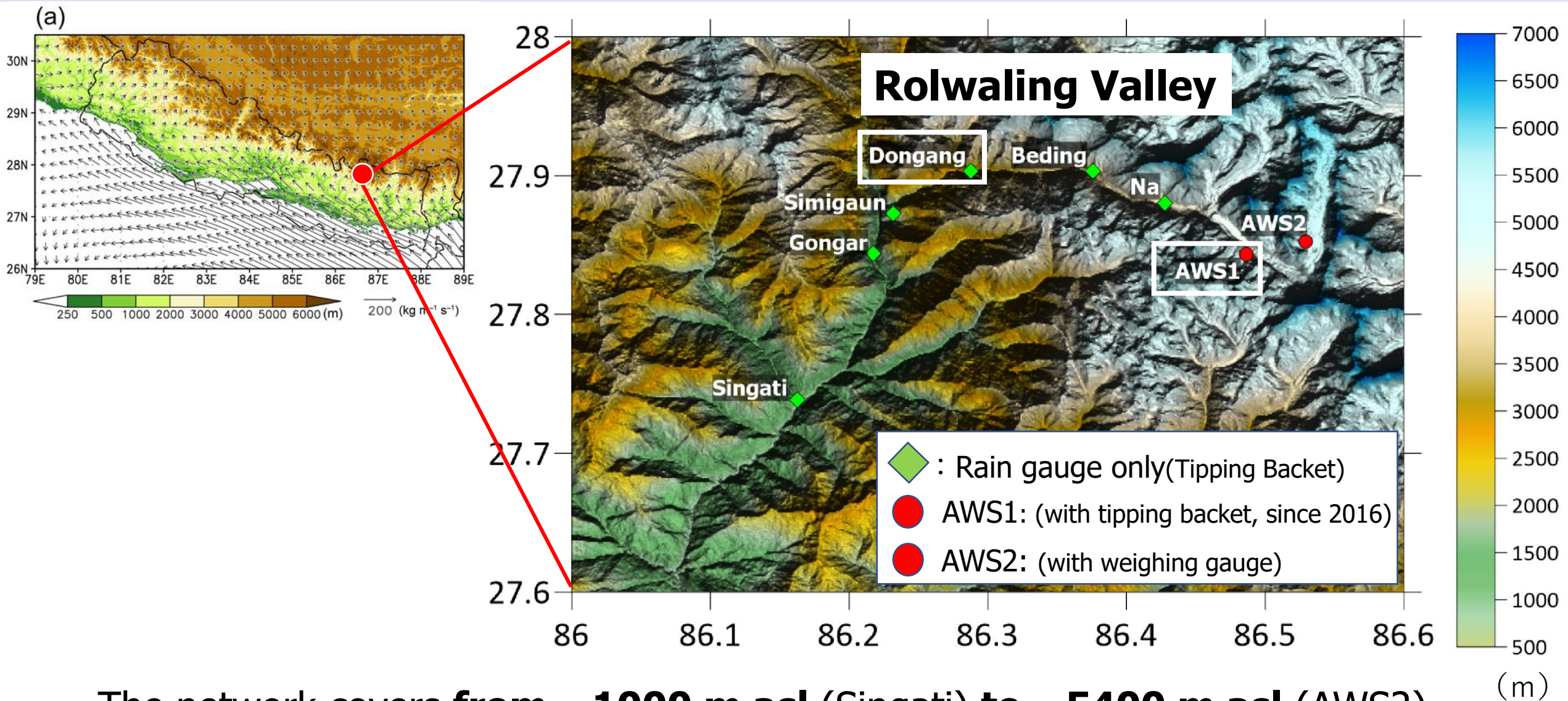
HiPRECS (under AsiaPEX):

- An international joint research project between **Japan** and **Nepal** funded by JSPS-KAKENHI (PI:H. Fujinami, Period: **Oct. 2018 – Mar. 2023**, with **KU, NAST and ICIMOD**)
- To understand **summer precipitation and its variability** in the higher elevations of the Himalayas (e.g., around Rolwaling glacier and Ponkar glacier)
- By ***In-situ* observation** (rain gauges and AWSs), **satellites** (GPM-DPR, GPM-GMI, CloudSat-CPR and Meteosat), **EAR5** reanalysis and **Cloud resolving regional models** (CReSS, WRF and NHM)



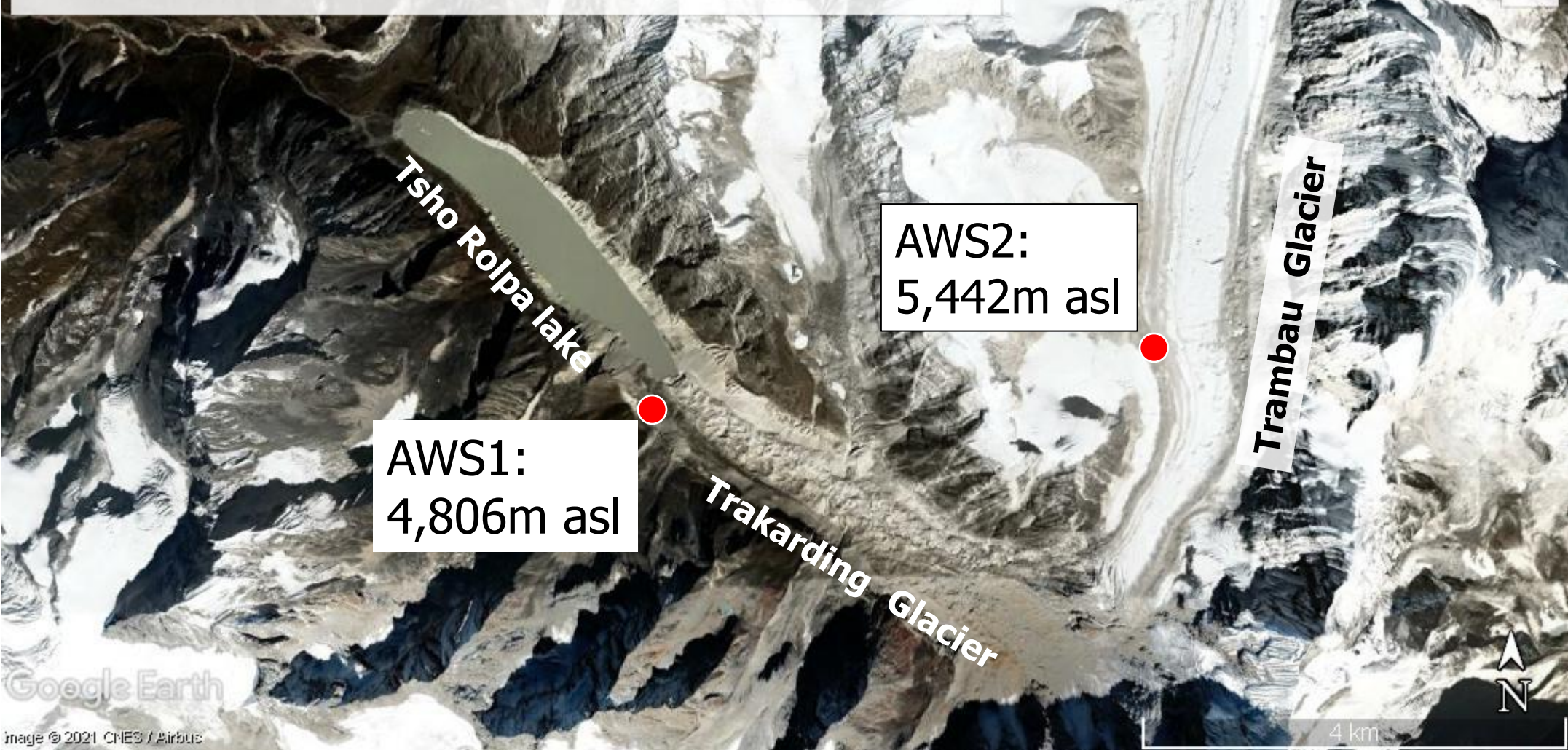
Rain gauge network along the Rolwaling Valley since 2019

in collaboration with JSPS-SNSF (PI:Prof. K. Fujita, Glaciologist)



The network covers **from ~1000 m asl (Singati) to ~5400 m asl (AWS2)**.

Tsho Rolpa glacier lake and Trakarding-Trambau glacier system

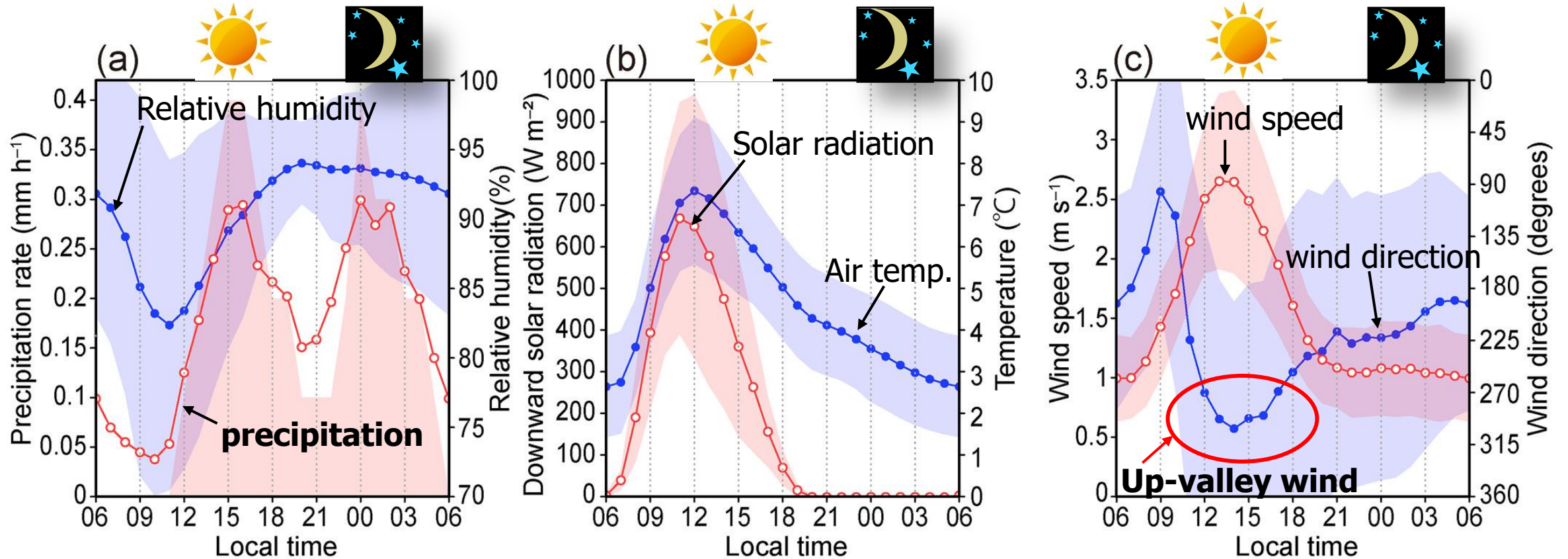


(©Google earth)

Diurnal cycle: Twice-daily maxima of precipitation

Observation at AWS1

3-year mean diurnal cycles (2016, 2017 and 2018)



(Fujinami et al. 2021, JGR atmos)

- **Daytime precipitation peak:**

Up-valley flows induced by surface heating causes condensation along the slope.

- **Nighttime precipitation peak:**

No signals in surface meteorological elements

➡ **Large-scale low-level monsoon flow**

Twice-daily maxima of precipitation at higher elevations based on in situ observations

★ Marshandi River Basin:

- Barros et al. (2000, *GRL*); (2001, *Natural Hazards and Earth System Sciences*)

▲ Pyramid station:

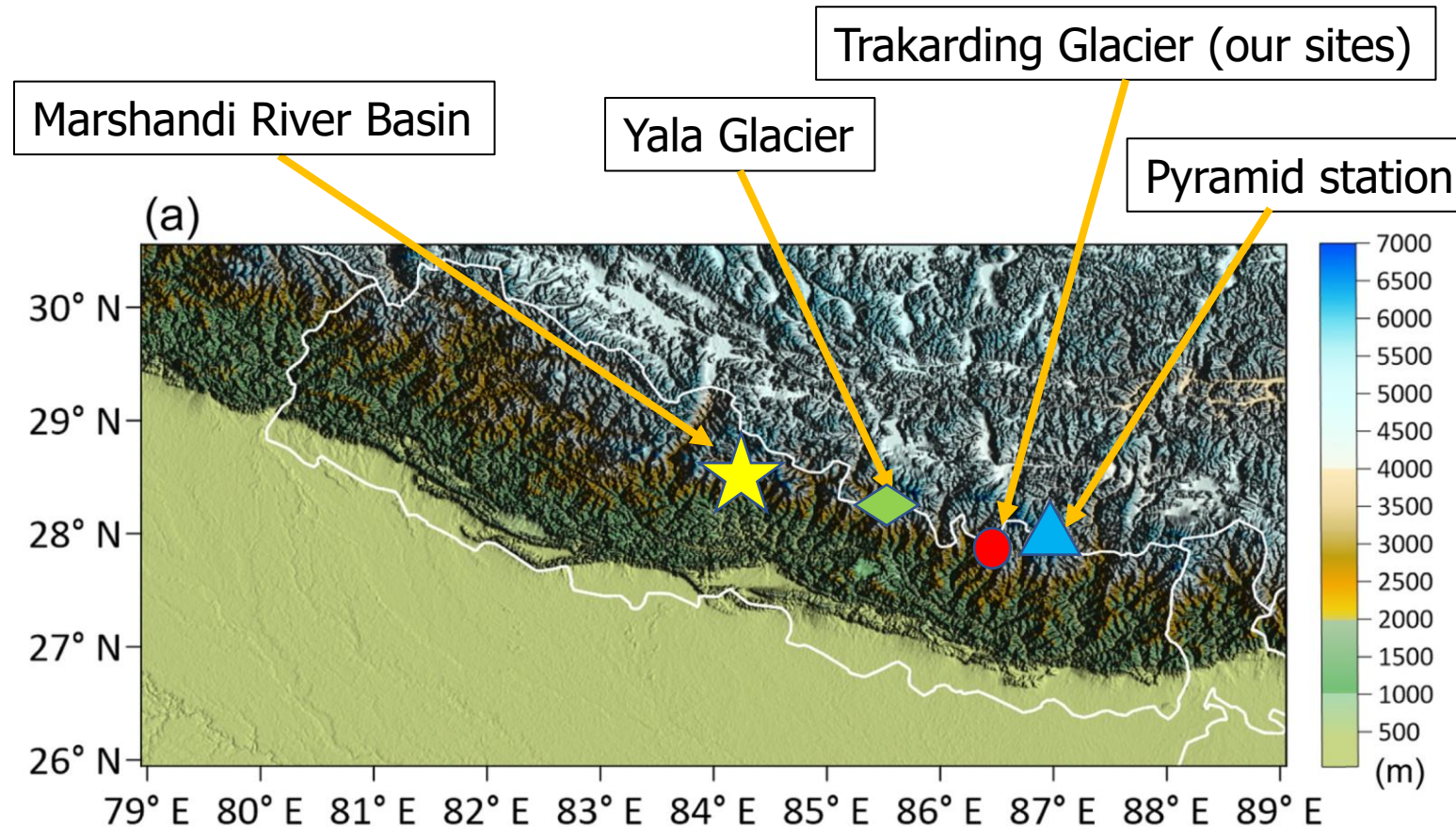
- Ueno et al. (2008, *MWR*), Yang et al. (2018, *IJOC*)

◇ Yala glacier:

- Shea et al. (2015, *International Journal of Water Resources Development*)

● Trakarding glacier:

- Fujinami et al. (2021, *JGR-Atmospheres*)

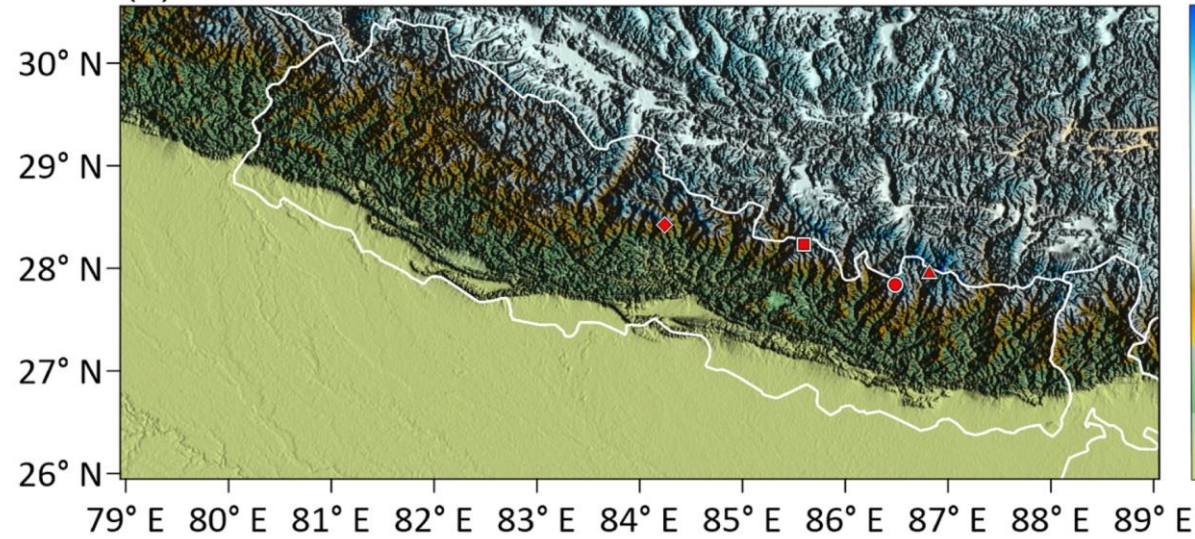


The similar diurnal cycle was reported in the past field experiments at the higher elevations.

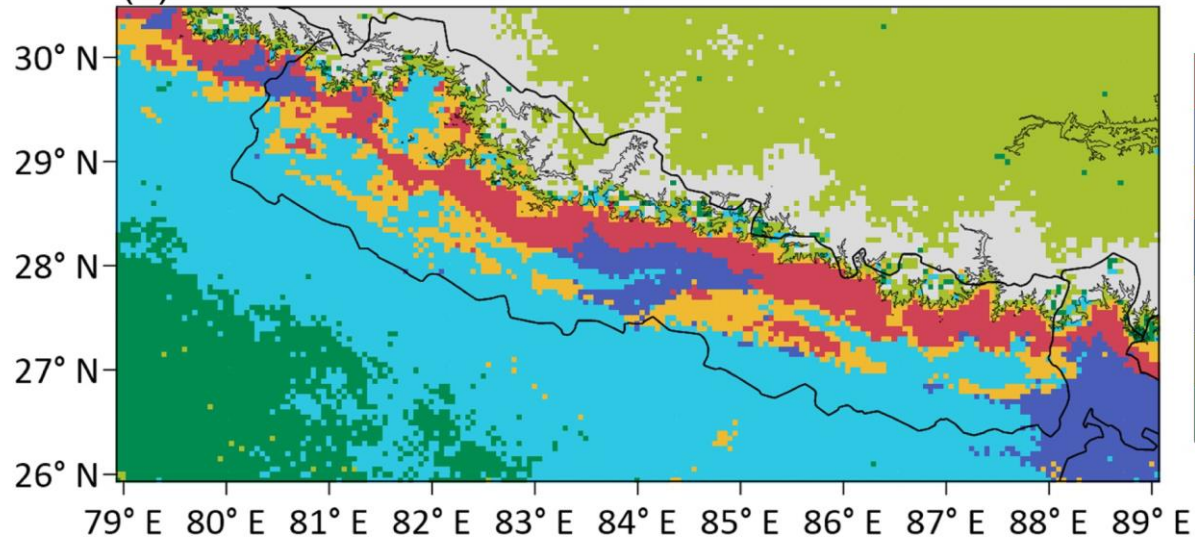
Are the **twice-daily maxima** unique to the high-elevation area?

Cluster analysis of diurnal cycle of rainfall frequency based on **TRMM-PR(1998–2014)**

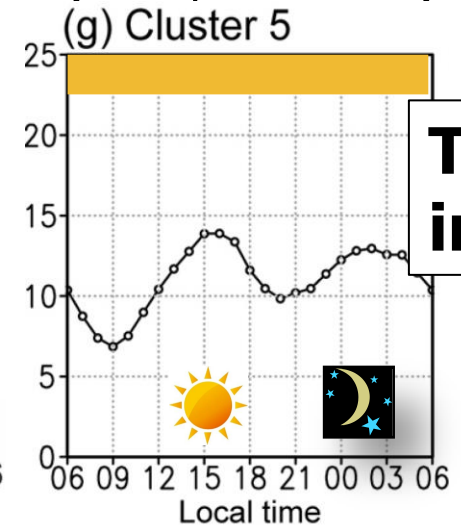
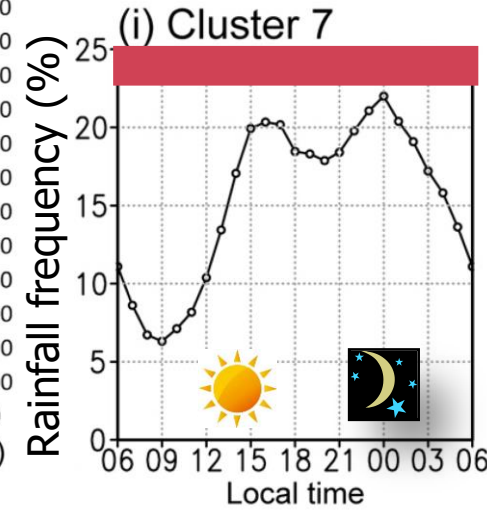
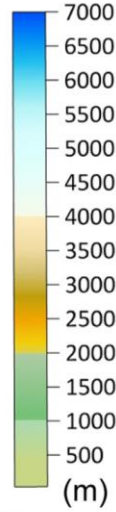
(a) Topography



(b) Distribution of clusters

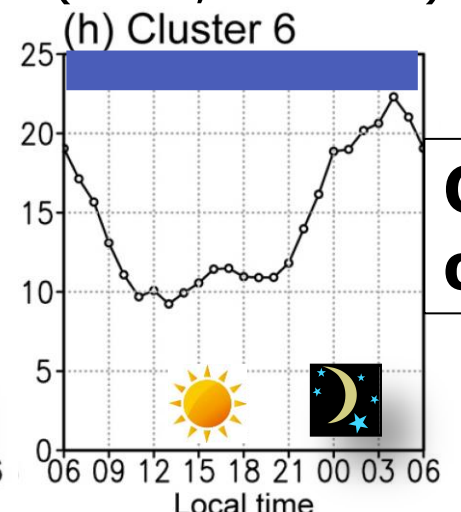
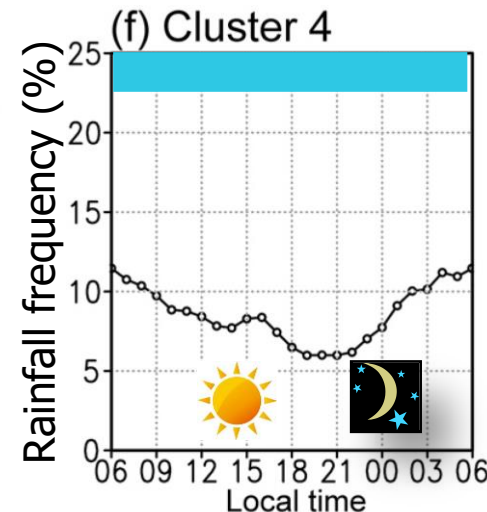
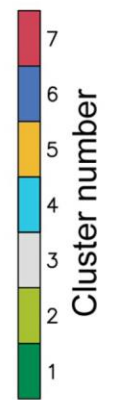


Higher elevations (>~1,500m asl): Cluster 5 & 7



Two peaks
in  & 

Lower elevations (<~1,500m asl): Cluster 4 & 6

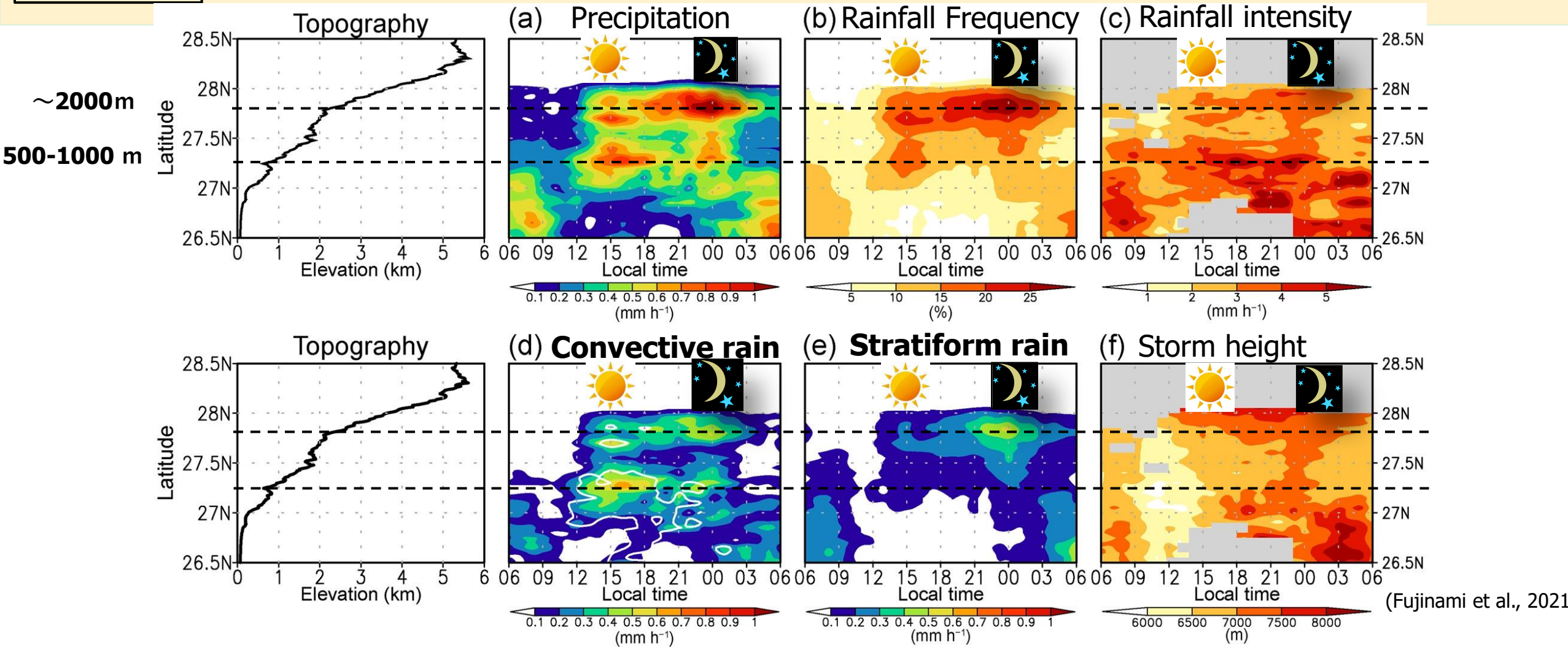


One peak
only in 

Contrasting precipitation properties in the two peaks

TRMM-PR

Latitude-local time section (85.5–86.5E)

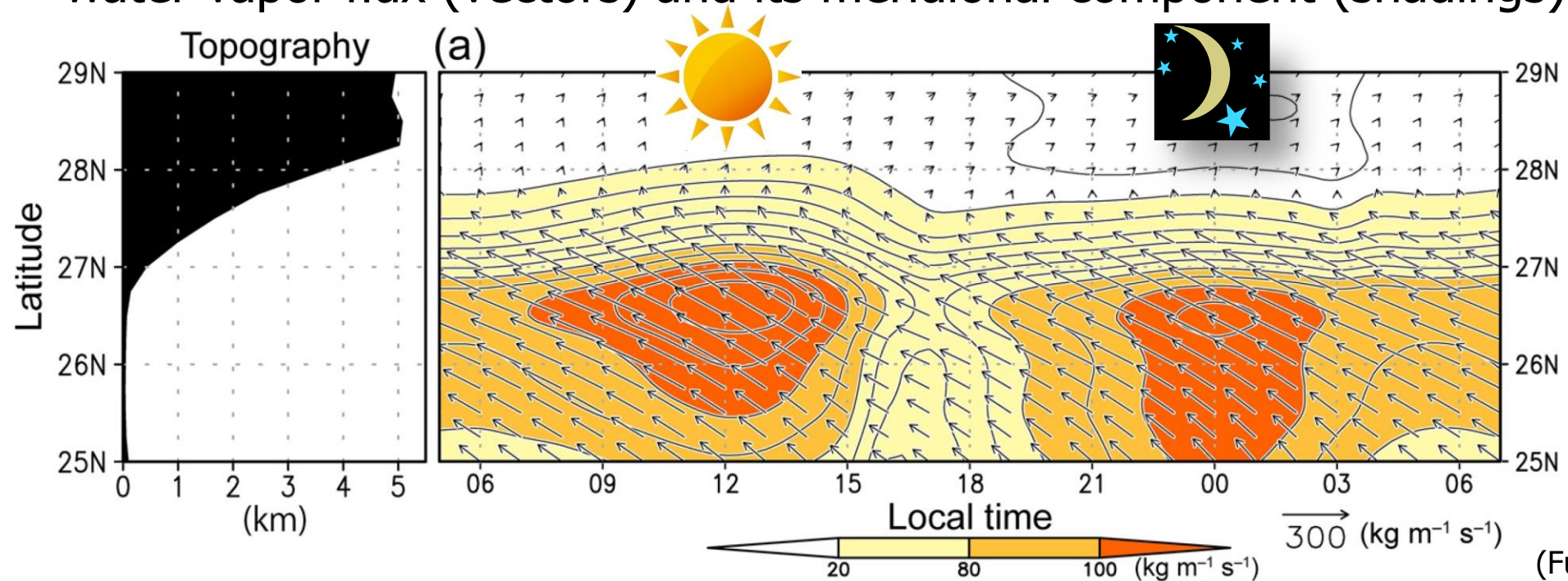


- **Daytime peak: convective-type** precipitation and **lower** rain-top height
- **Nighttime peak: stratiform-type** precipitation and **higher** rain-top height

Mechanism for driving the twice daily maxima of precipitation

ERA5

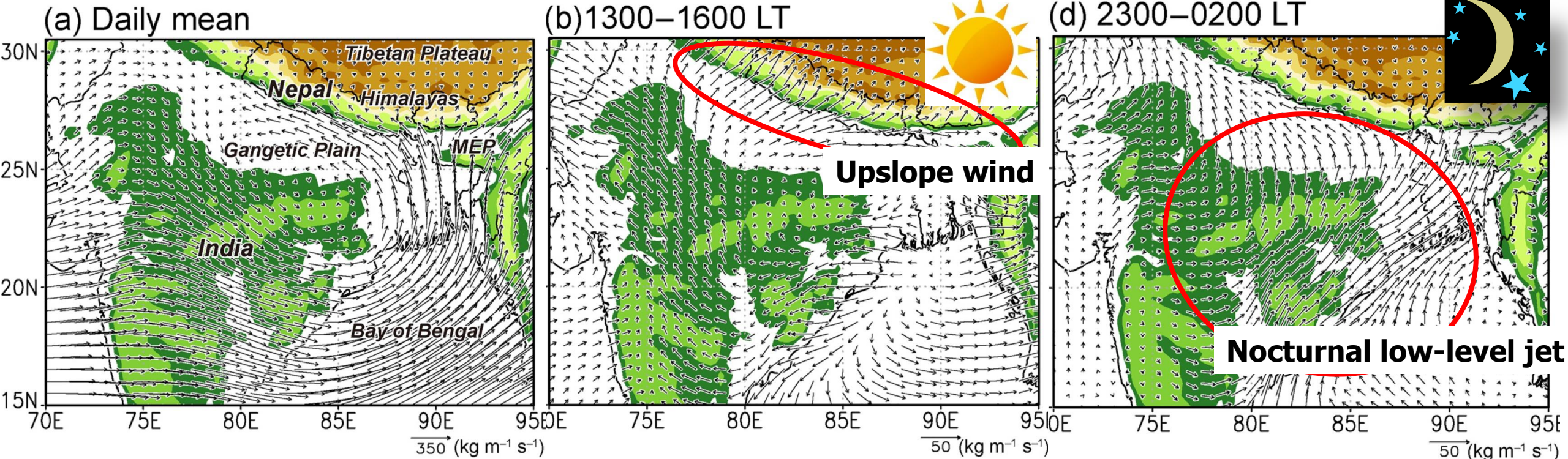
Latitude-local time cross section (85.5–86.5E)
water vapor flux (vectors) and its meridional component (shadings)



(Fujinami et al. 2021)

- The moisture flux strengthens twice a day during the day (12:00) and night (00:00).
- The timing corresponds to the twice-daily maxima of precipitation.

Contrasting pattern of water vapor flux anomalies



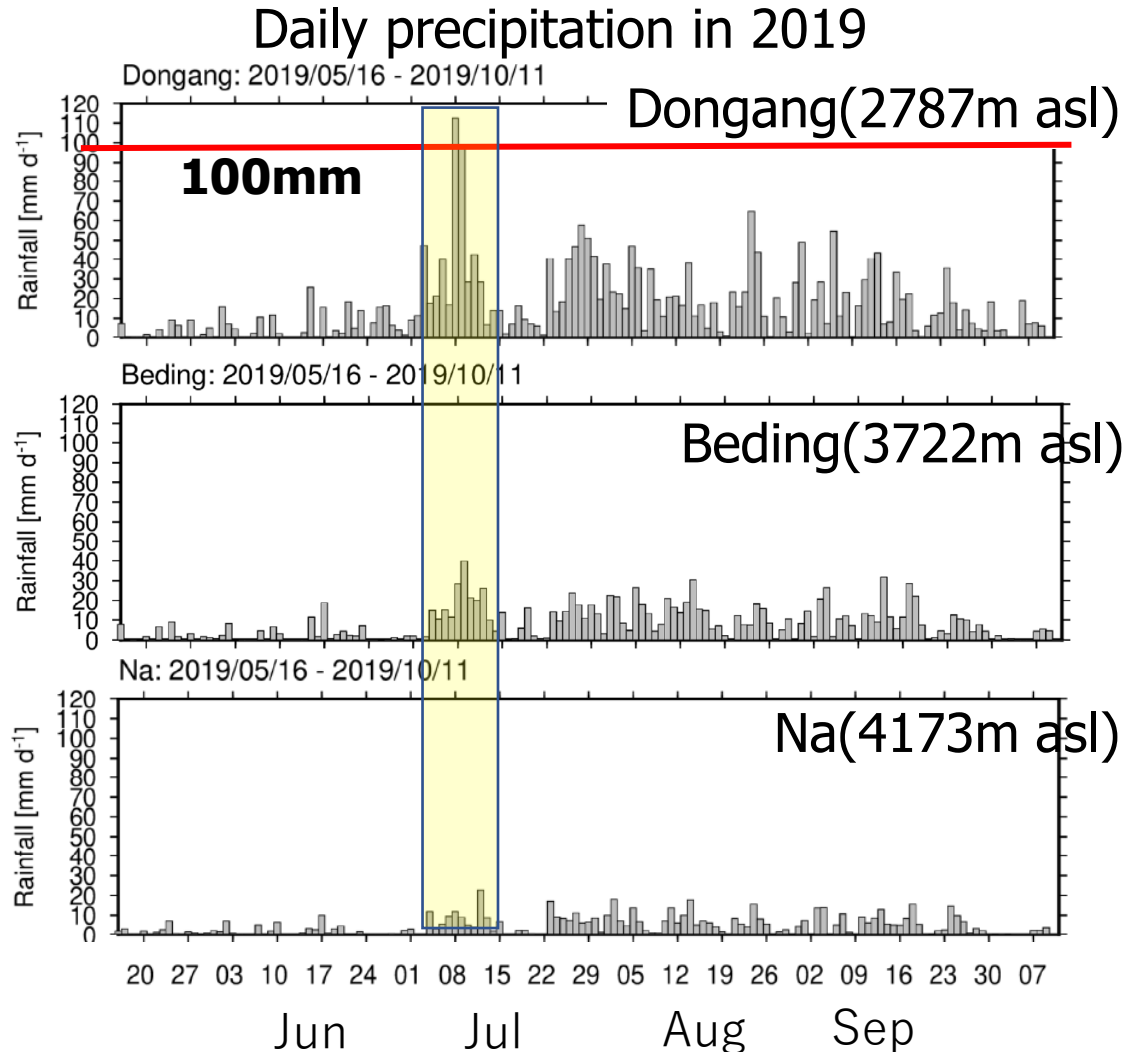
(Fujinami et al. 2021)

- **Daytime:** Southerly moisture flux anomalies (upslope flow) around the Himalayas; Strong upslope flow due to surface heating -> **convective-type rain**
- **Nighttime :** Large-scale moisture flux anomalies toward the southern slopes; Weak upslope flow and large-scale moisture flux convergence -> **stratiform-type rain**

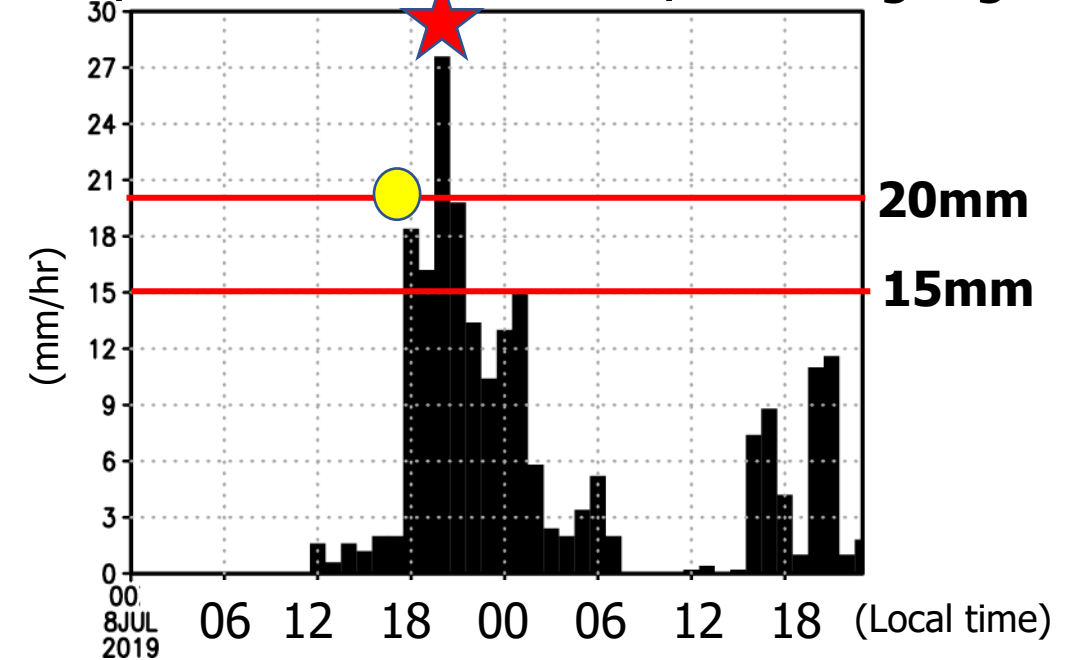
A heavy rainfall event due to a monsoon low through multi-scale process

Observation at Donggang (2,800m asl)

A heavy rainfall event on 8–9 July 2019



Hourly rain rate on 8–9 July at Donggang

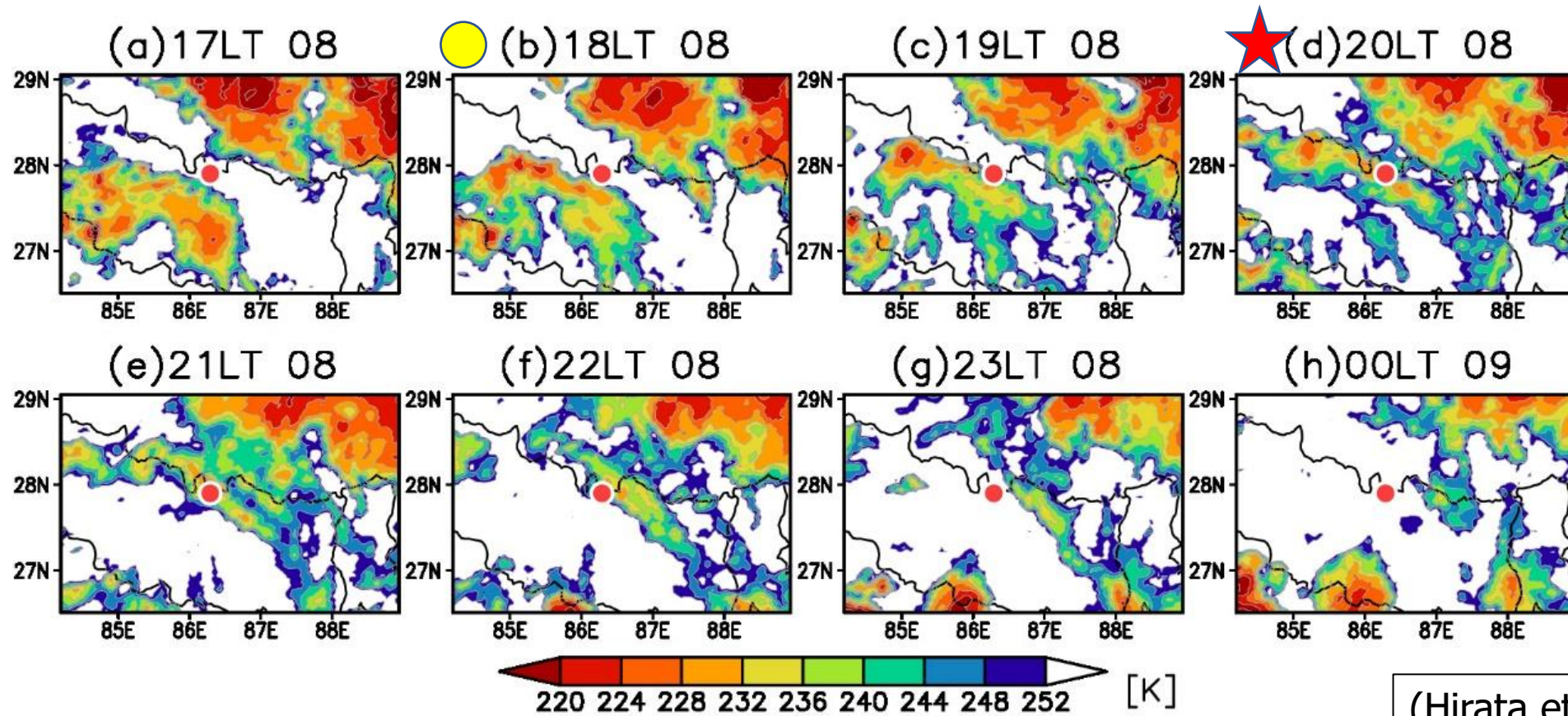


- High precipitation at Donggang (>200mm) and Beding (>60mm) during 8–9 July
- High rain intensity during 20:00–21:00 local time (>20mm/hr)

Cloud system during the heavy rainfall event

METEOSAT IODC

Hourly IR images from METEOSAT IODC

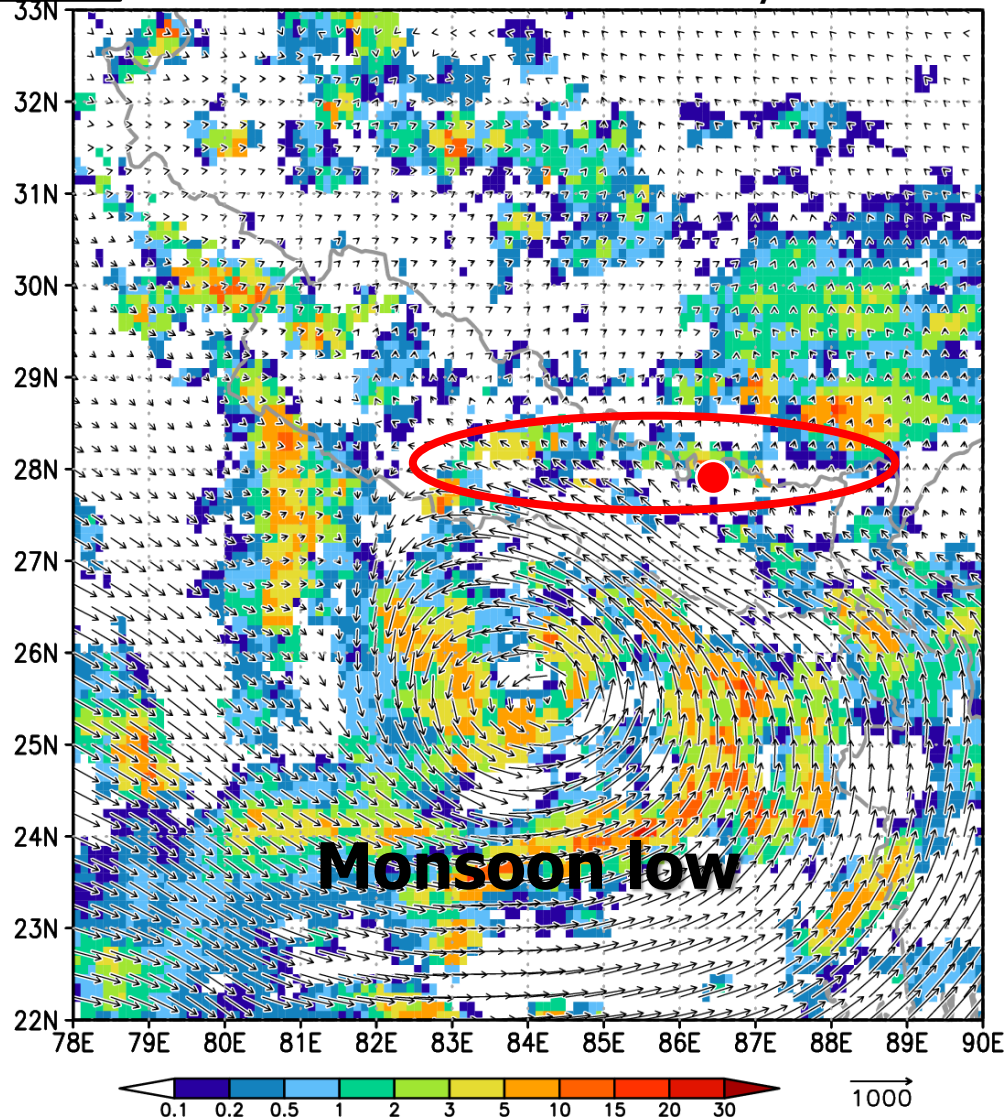


- **Northward moving line-shaped cloud** band with **meso-scale systems**
- The cloud band passed over Donggang during the heavy rainfall event.

Synoptic-scale atmospheric circulation during the heavy rainfall event

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At 21:00 LST on 8th July

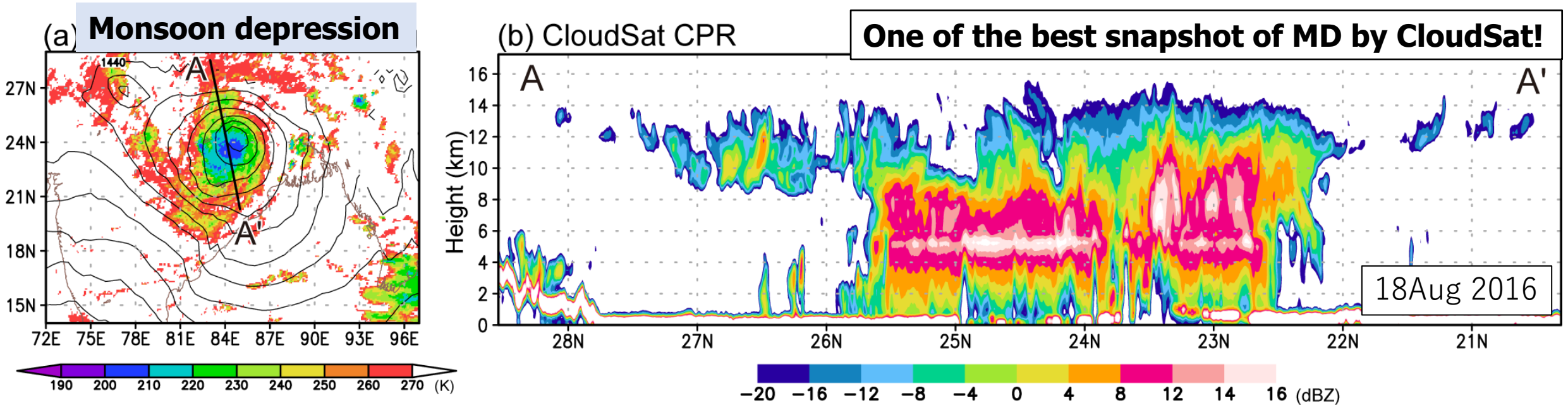


- A **monsoon low** by the Himalayas
- Large **southeasterly moisture transport** toward the Himalayan slopes
- The **Himalayan slopes** and the southeasterlies appear to cause the cloud/rain band.

Shadings: Rainfall (**IMERG**,mm/hr) 21:00 NPT 8July 2019,
Vectors: vertically integrated water vapor flux (**ERA5**)

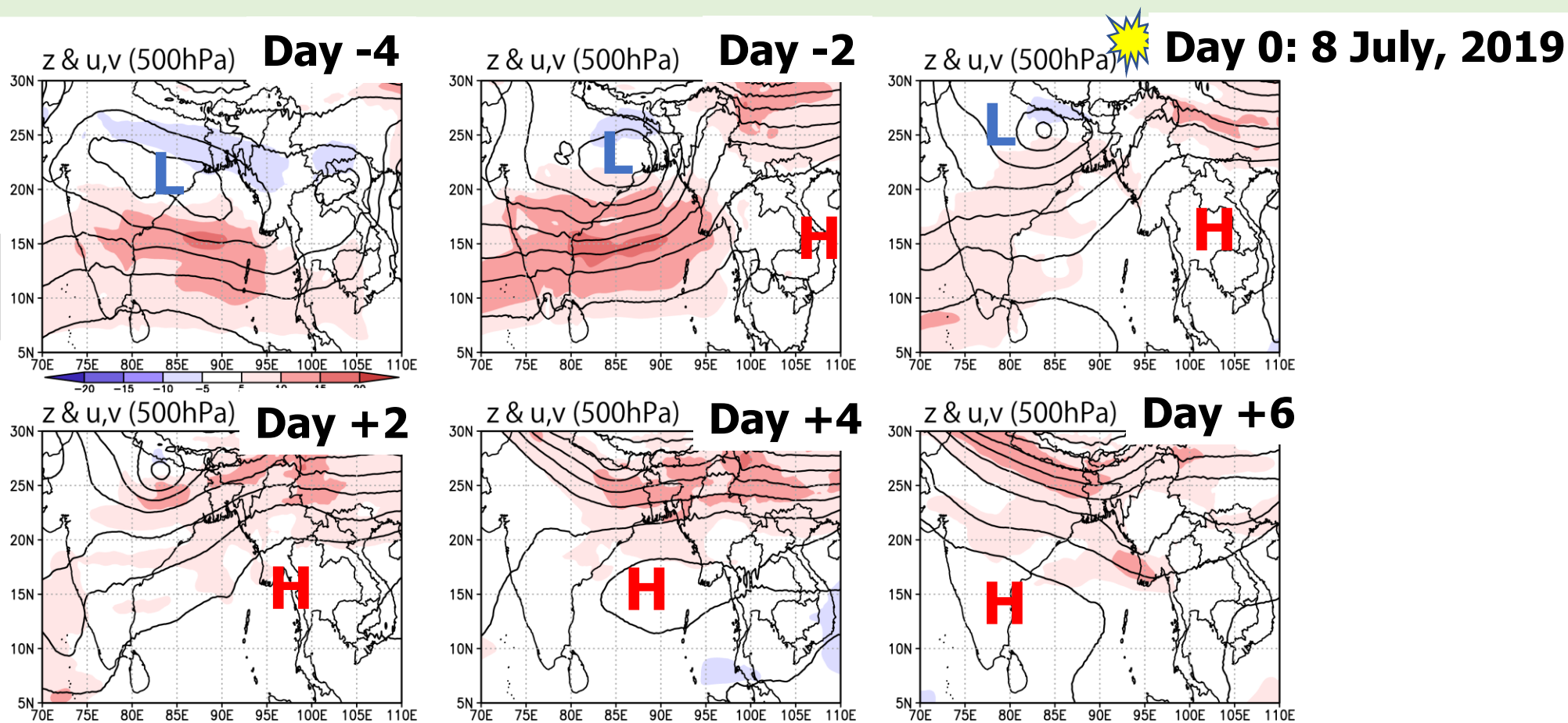
Monsoon lows (low pressure system: LPS)

- **Effects of monsoon lows** (synoptic-scale disturbances over South Asia):
 - Onset of rainy season (e.g., Barros and Lang 2003) and extreme rainfall (e.g., Bohlinger et al. 2017)
 - to **moisten the (middle) troposphere** to the south of the Himalayas
 - to **enhance water vapor transport** toward the southern slopes
- **Locations of monsoon lows are spatially clustered by ISOs (BSISO and QBW).**
(e.g. Goswami et al. 2003; Hatsuzuka & Fujinami 2017)



(Fujinami et al. 2020, QJRMS)

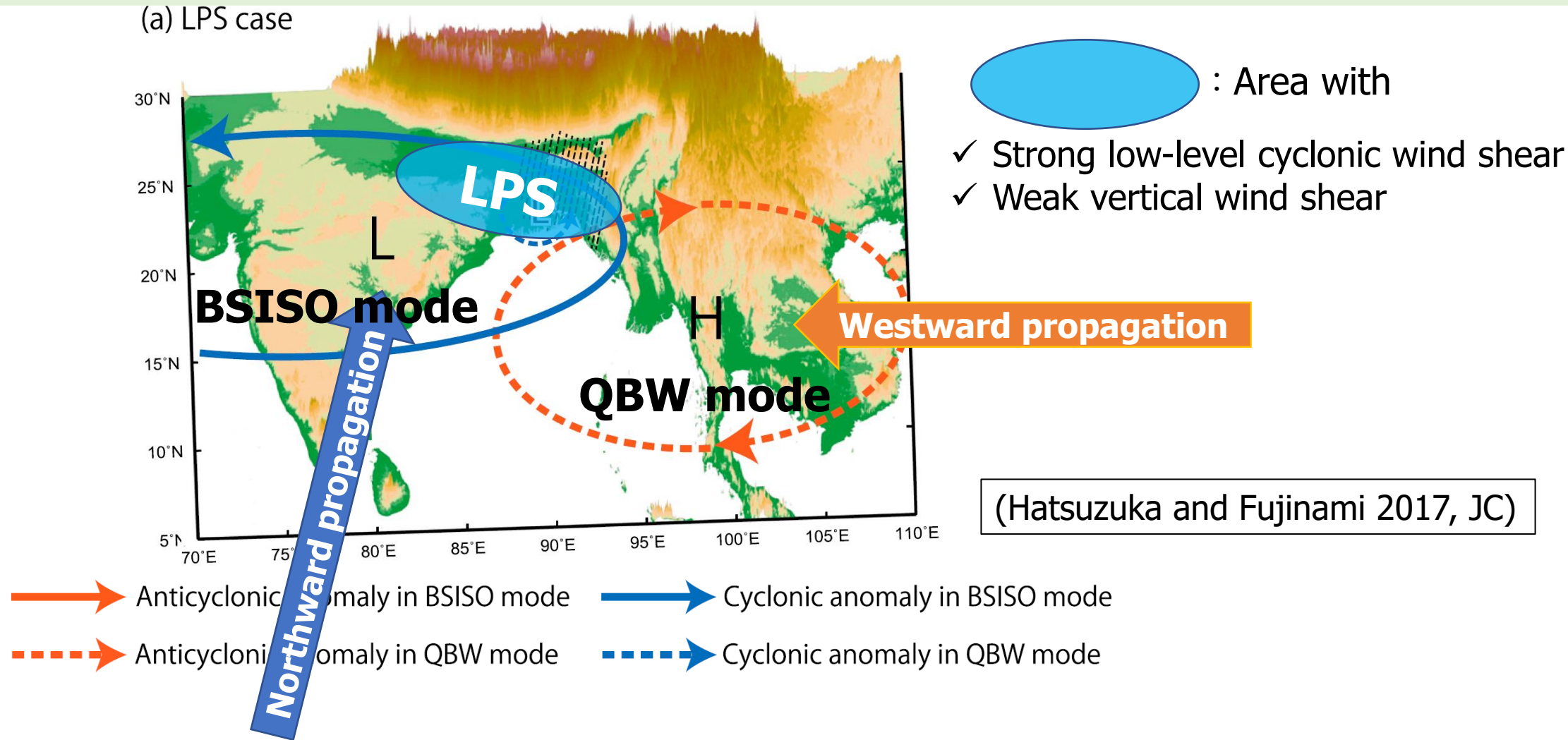
Time sequence of geopotential height and zonal winds at 500 hPa



- Westerlies move northward toward the Himalayas from day -4 to day 0.
- **Westward propagating high** over the BoB push them northward (**QBW mode**).

ISOs → low pressure system → Line-shaped cloud/rain band → Mesoscale convection

Relationship between ISOs (BSISO and QBW) and LPS over the Gangetic Plain



- Enhancement of the low-level **cyclonic wind shear** over the Gangetic Plain
- A favorable environment for the genesis and the track of the LPS

Summary

In situ observation in the higher elevations of the Himalayas:


- Meteorological observations at elevations of $>4,000$ m asl over the Himalayas are still extremely scarce. They are essential to understand the **high mountain water cycle** with **glaciers (HiPRECS)**.

Diurnal cycle of precipitation:

- **Twice-daily precipitation maxima**, corresponding to day and night, occur widely across the higher elevations of Nepal Himalayas, including in the glacierized area.
- Upslope flows due to surface heating drive the daytime rainfall peak, and a monsoon nocturnal low-level jet creates the nighttime peak (**Land surface effects**).

The heavy rainfall event due to the multi-scale process:

- **Multi-scale process** (ISOs, LPS, line-shaped rainband and meso-scale convection) led to the heavy rainfall event at Dongang (2,800 m) on 8–9 July, 2019.
- Understanding of multi-scale processes from the tropics to the mid-latitudes are needed to better understand precipitation variability in the Himalayas.



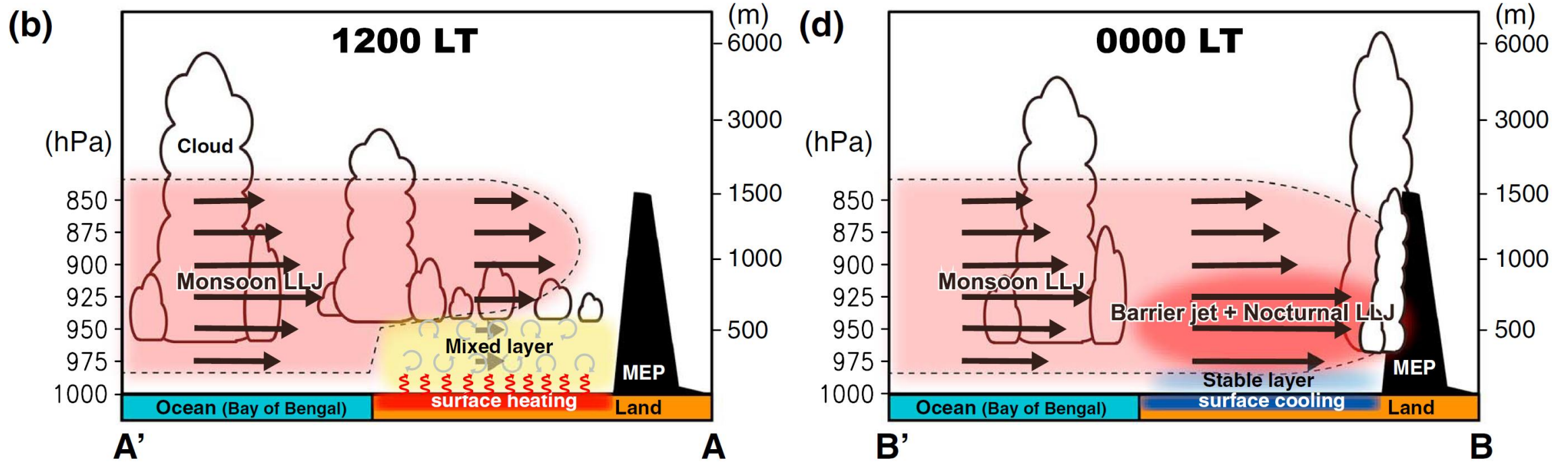
Thank you for your attention

@Karche in Rolwaling

Status of data acquisition in summer precipitation (May–October)

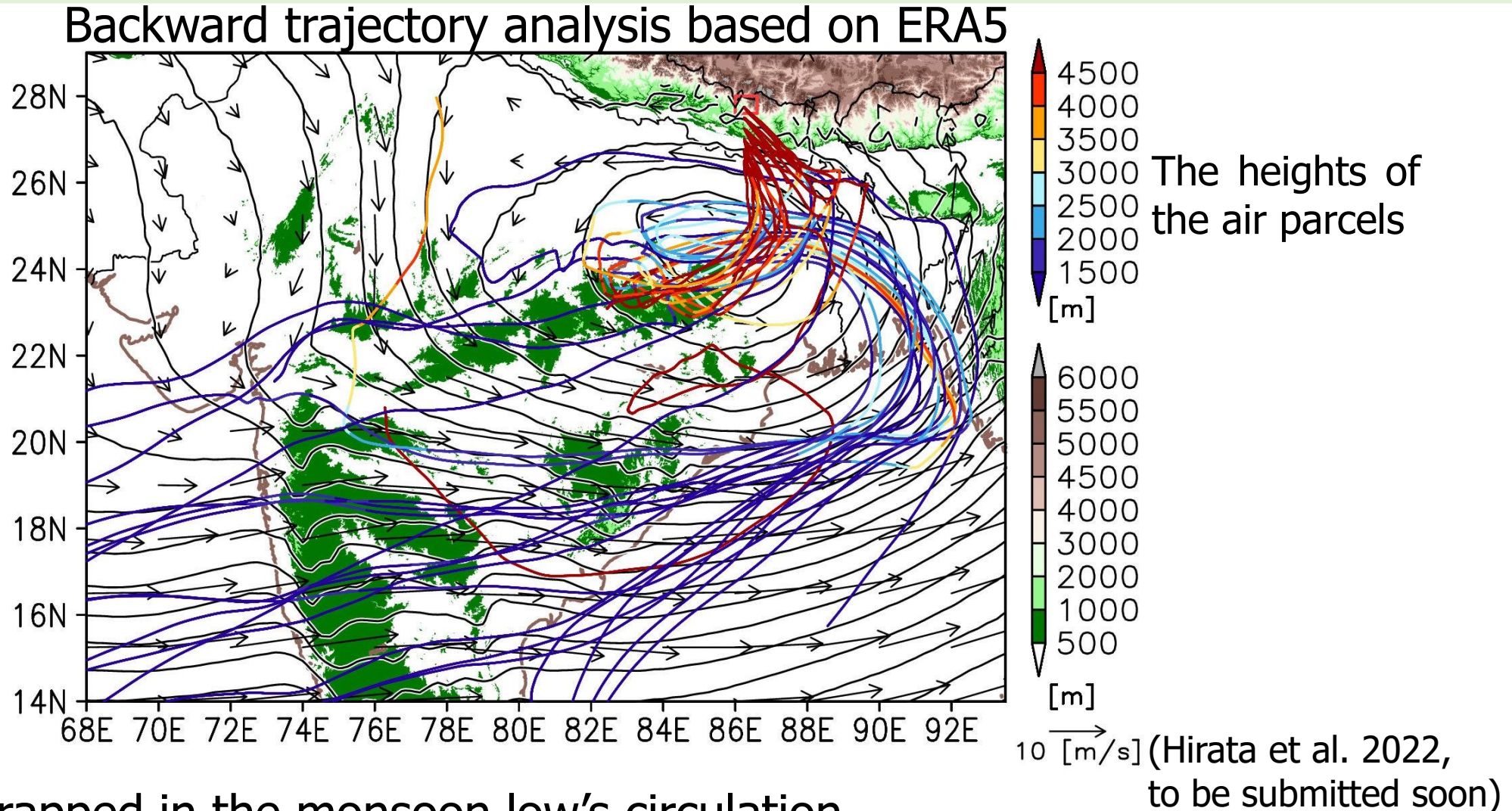
	2016	2017	2018	2019	2020	2021
Singati				○	○	○
Gongar				○	○	○
Simigaun				○	○	○
Dongang				○	○	○
Beding				○	○	○
Na				○	○	○
AWS1	○	○	○	N/A	○	○
AWS2					N/A	N/A

Effects of ABL process on the low-level horizontal winds



Fujinami et al. (2017, JGR Atmos)

How does the monsoon low transport the humid air into our observation site during the heavy rain event?



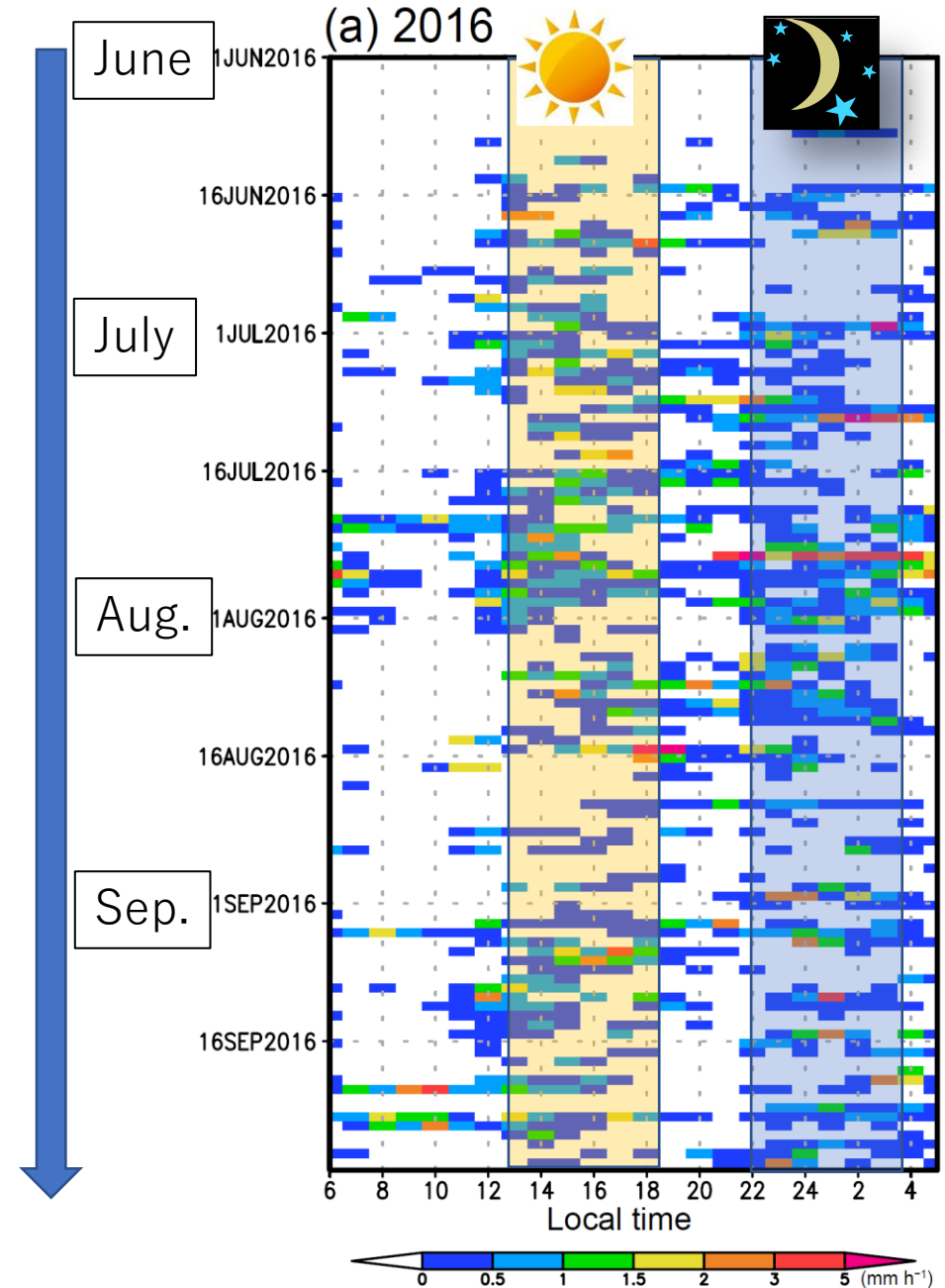
- Air parcels trapped in the monsoon low's circulation
- **Moisture transport** into the observation sites in the **middle troposphere**

Results

In-situ observations

- Precipitation in **daytime** and **nighttime**
- **Weak rainfall rate:**
Precipitation events with rainfall rates of $\leq 0.5 \text{ mm h}^{-1}$ account for $\sim 60\text{--}70\%$ of all precipitation events.

Season-local time cross section



Data

- **Automatic Weather Station (AWS) at 4,806 m asl by Trakarding glacier in Rolwaling region, eastern Nepal (Fujita et al. 2021):**

- **Hourly data, 2016–2018, June–September(JJAS)**
- precipitation, air temperature, wind speed and direction, relative humidity and solar radiation
- Precipitation: tipping bucket type with 0.1 mm resolution

- **TRMM-Precipitation Radar (PR) 2A25 v7:**

- **Hourly** diurnal climatology (**1998–2014**, JJAS)
- Near-surface rain, rain types and storm height
- **0.05 × 0.05 grid**

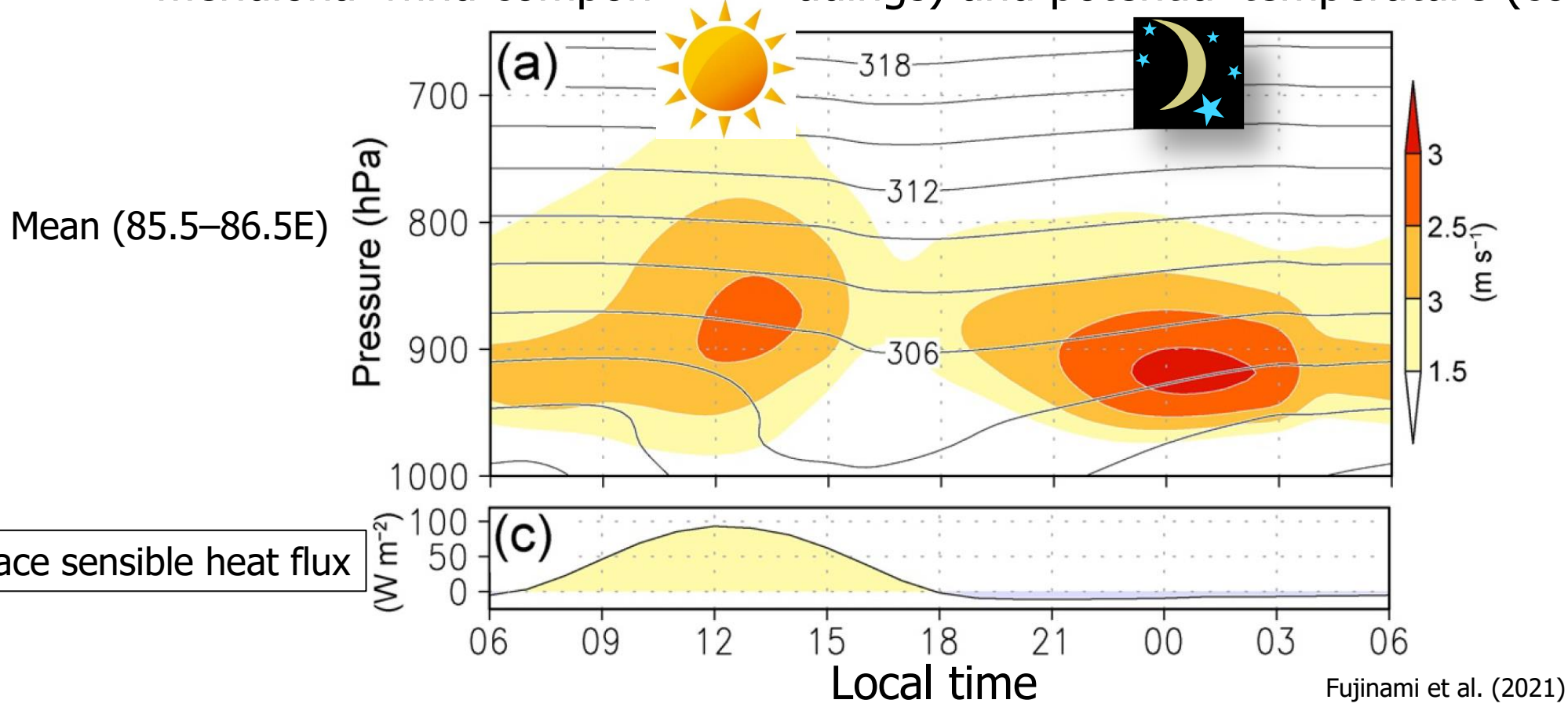
- **ERA5 reanalysis:**

- **Hourly** diurnal climatology (**1998–2014**, JJAS)
- Meteorological elements: z, u, v, q and t
- **0.25 × 0.25 grid**

Diurnal cycle of horizontal winds due to the development of the ABL

ERA5

Vertical-local time cross-section over the windward plain:
meridional wind component (shadings) and potential temperature (contours)

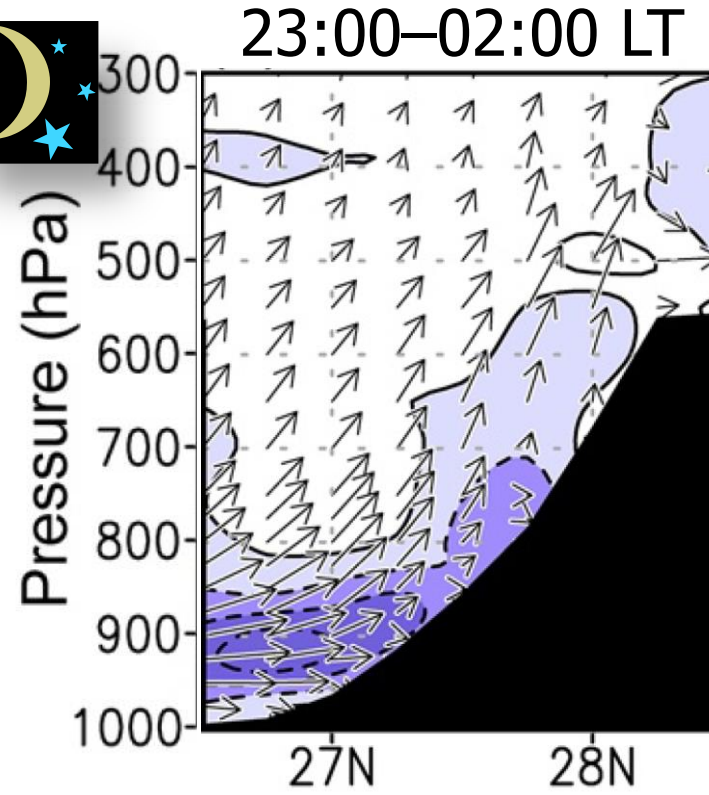
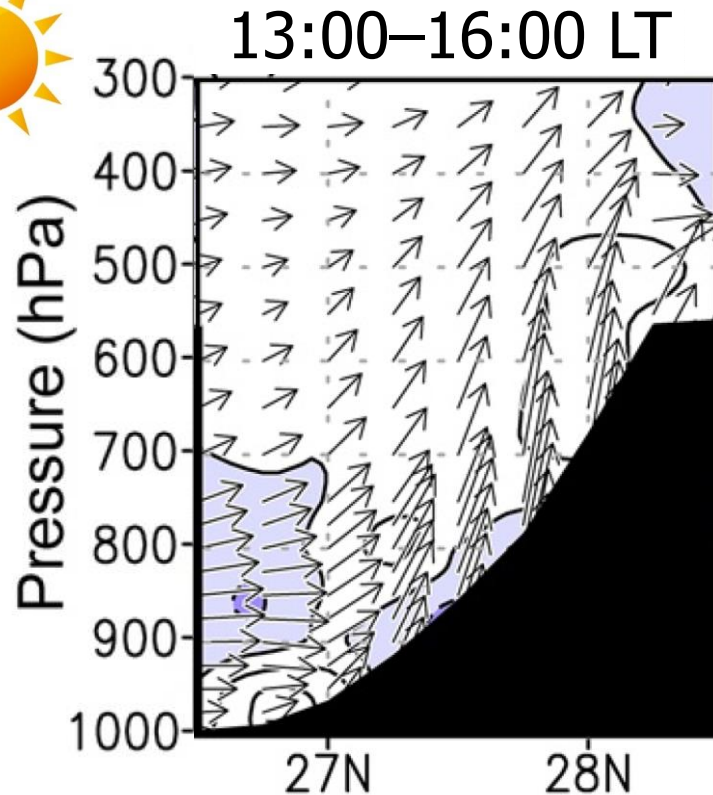


- **Daytime:** Upslope flow over the slopes (**Anabatic winds over the slopes**)
- **Nighttime :** Acceleration of horizontal air flow (**Nocturnal jet over the plain**).

Contrasting vertical structure of meridional circulation

Latitude-height cross-sections of meridional and vertical winds (vectors) and horizontal water vapor flux divergence (shadings)

ERA5



Fujinami et al. (2021)

- **Daytime:** Strong **upslope flow along the slope surface** due to surface heating
- **Nighttime :** Weak upslope flow, the enhancement of southerlies below 850 hPa, **large-scale moisture flux convergence.**

Day lag composite of 850-hPa streamfunction, wind and OLR anomalies in day -4, -2 and 0

Quasi-biweekly (QBW) mode

Equatorial Rossby wave (n=1) as a key agent that regulates the low-level zonal wind on QBW time-scales

- Westward propagating double vortices
- Westward phase speed: ~ 6.5 m/s
- Zonal wavelength: $\sim 6000-7000$ Km
- The period of ERW ranges submonthly time-scales by nature

(Fujinami et al. 2014 *Clim. Dyn.*)

